







DIM ESEE-2 innovative workshop

DIM ESEE 2021: Innovation in exploration

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Shallow geophysical investigations by combining seismic, geoelectric and direct-push logging methods

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Course Description

- Shallow geophysical methods in brief: seismic refraction, multielectrode direct current geoelectric method, direct-push logging (borehole engineering sounding) methods
- Environmental investigations made by combining seismic, geoelectric and direct-push logging methods
- The workflow of inverse modeling methods
- The problem of estimating the layer boundaries in joint inversion of surface geophysical data
- Joint inversion of different direct-push logging datasets
- Evolutionary meta-algorithmic inversion method. Estimating the petrophysical and zone parameters of near-surface structures as input for soil mapping and mining (e.g., gravel) applications









DC Geoelectric Method



(Wenner array: MN=AB/3)

Depth of investigation \leq AB/3

Apparent Resistivity:
$$\rho_a = k \frac{\Delta U}{I} = k \frac{\Delta U_{MN}}{I_{AB}}$$









Multielectrode (ERT) method

Dipole-dipole array

(Reference depth = $R_{(AB/2,MN/2)}/2$)







ERT Survey Along Riverbank

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ERT Survey for Rock Differentiation



NUMBER OF READINGS: 362

SPACING BETWEEN ELECTRODES: 5 m

SITE: Spain

GEOLOGY: granitic area with altered and fractured zones

DURATION OF MEASUREMENTS: 1 hour (system set-up + reading taking)



hard granit: more than 2 000 ohm.m







Direct Push (Logging) Method

Engineering Geophysical Sounding



- 1 Vehicle
- 2 Semitrailer
- 3 Hydraulic machinery
- 4 Pressure piston
- 5 Measuring tube
- 6 Measuring head
- 7 Anchor

Fejes and Jósa (1990)

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Characterization of Soils

2@21

ĒSEĒ

MEASURED SOUNDING PARAMETERS



pore volumen

content

pore volumen









Direct Push NMR Logging Silty unsaturated zone

Level of water table



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Seismic Method



- a) Vertical reflection
- b) Wide angle reflection

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- c) Critical refraction
- d) Direct wave











Workflow of Inverse Modeling







1D Inversion of Vertical Electric Sounding Data

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H type: $\rho_1 > \rho_2 < \rho_3$ Q type: $\rho_1 > \rho_2 > \rho_3$ K type: $\rho_1 < \rho_2 > \rho_3$ A type: $\rho_1 < \rho_2 < \rho_3$ Supported by





- We assume that the layers are homogeneous and isotropic
- Estimated parameters of the 1D model are:
 - h_i: layer thicknesses and
 - ρ_i : layer (true) resistivities









2D Inversion of Geoelctric Datasets









1.5D Inversion of Geoelctric Datasets



Gyulai et al. (2014)

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- Variations of layer boundary and resistivity along the profile are described by continuous functions **Joint inversion scheme**
- Discretization of layer parameters is based on series expansion (Dobróka, 1993)

$$m_i(x) = \sum_{q=1}^{Q^{(i)}} B_q^{(i)} \Phi_q(x)$$

where m_i denotes the i-th physical or structural parameter, B_q is the q-th expansion coefficient, Φ_q is the q-th basis function (up to Q number of additive terms)

• Number of unknowns is

- Set of 1D local inversions = 105

- Series expansion-based 2D inversion = 9

- Highly overdetermined inverse problem is solved for the expansion coefficients (the basis functions are known quantities)
- Higher accuracy, reliability, resolution and stability (no smoothness constraints)











Joint Inversion of Geoelctric Datasets







Inversion of Refraction Data







80



Inversion for Identical Layer Boundaries





Results of geoelectric-seismic joint inversion



Results of geoelectric 1.5D inversion



Results of geoelectric-seismic joint inversion











Inversion for Different Layer Boundaries



ACADEMY

Tamás Ormos (2006)







Nuclear Waste Site Characterization









Genetic Inversion of Direct Push Data

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- Forward problem (model >> predicted data)
- Zone parameters in the shallow unsaturated zone
- Fitness function of the 1D inverse problem

 $d_{s}^{(\text{calculated})} = g_{s}(V_{w}, V_{cl}, V_{ma}, c)$

$$\mathbf{c} = \left[GR_{cl}, GR_{ma}, R_{cl}, \rho_{cl}, \rho_{ma}, \Phi_{N,cl} \right]^{T}$$

$$F(\mathbf{c}^{(j)}) = -\left[(NS)^{-1} \sum_{n=1}^{N} \sum_{s=1}^{S} \left(\frac{D_{ns}^{(m)} - g_s(\mathbf{m}^{(n)}, \mathbf{c}^{(j)})}{D_{ns}^{(m)}} \right)^2 \right]^{1/2}$$

$$\mathbf{F}^{*}(\mathbf{c}^{(j)}) = -\left[\left(\mathbf{N}^{*}\mathbf{S} \right)^{-1} \sum_{h=1}^{H} \sum_{n=1}^{N_{h}} \sum_{s=1}^{S} \left(\frac{\mathbf{D}_{hns}^{(m)} - \mathbf{g}_{s}(\mathbf{m}^{(hn)}, \mathbf{c}^{(j)})}{\mathbf{D}_{hns}^{(m)}} \right)^{2} \right]^{1/2}$$













Genetic Algorithm Based Search







Exchange in position j=2









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QC of Inversion Results



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QC of Inversion Results













2D Genetic Meta-Algorithmic Inversion





















2D Genetic Meta-Algorithmic Inversion









Hyperparameter Estimation for Zone Parameters

Zone parameter	Search domain	Estimated Value	Estimation Error	Unit
GR (clay)	8.0-12.0	8.02 (<mark>8.86</mark>)	0.17 (<mark>0.01</mark>)	kcpm
GR (matrix)	0-2.0	1.98 (<mark>1.83</mark>)	0.01 (<mark>0.01</mark>)	kcpm
ρ (clay)	1.9-2.3	1.97 (<mark>2.07</mark>)	0.03 (<mark>0.02</mark>)	g/cm ³
ρ (matrix)	2.3-2.7	2.31 (<mark>2.41</mark>)	0.01 (<mark>0.01</mark>)	g/cm ³
Φ _N (clay)	0.2-0.5	0.33 (<mark>0.43</mark>)	0.01 (<mark>0.05</mark>)	v/v
R (clay)	1.0-6.0	4.19 (<mark>4.56</mark>)	0.49 (<mark>0.61</mark>)	ohmm

*1D meta-algorithmic inversion *2D meta-algorithmic inversion Szabó N. P., 2018. A genetic meta-algorithm-assisted inversion approach: hydrogeological study for the determination of volumetric rock properties and matrix and fluid parameters in unsaturated formations. **Hydrogeology Journal** 26, 1935-1946, D1 ranked.

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Quadratic mean of standard deviation = 0.038 v/v



Quality Check of Inversion Results



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Thank you for your attention.

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